Supporting Online Material for

2500 Years of European Climate Variability and Human Susceptibility

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Published 13 January 2011 on Science Express
DOI: 10.1126/science.1197175

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Supporting Online Material for

2500 Years of European Climate Variability and Human Susceptibility

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Materials and methods

**Oak data.** Annually resolved oak ring width measurement series from three regions (Northeast France = NEF; Northeast Germany = NEG; Southeast Germany = SEG; Table S1) in Central Europe (CE) were compiled to continuously span the past ~2500 years and cover a large fraction of temperate forest area north of the Alpine arc and south of the Baltic Sea. The material contains oak (*Quercus robur* L. and *Q. petraea* (Matt.) Liebl.) wood from archaeological, sub-fossil, and historical surveys, as well as from recent findings. The CE oak species, *Q. robur* L. and *Q. petraea* (Matt.) Liebl., are not anatomically distinguishable (1). The historical portion of the oak dataset, which clearly represents the majority of the samples (Table S1), reaches continuously from the Iron Age to the early 20th century. We extended this dataset into the 21st century following a new approach to overcome the ‘update desideratum’ of modern site bias: adaptation of the recent to the historical data that was recently introduced by (2). In this case, recent oak beams and timber were randomly sampled at different sawmills and lumberyards scattered over the same area from which the historical
wood were derived. This random sampling strategy lowers site control and ecological understanding of the recent material, and typically results in artificial signal-degradation throughout the modern calibration period. It therefore helps to balance uncertainty-levels over multi-centennial to millennial-long tree-ring records composed of recent and historical wood material, and prevents from statistical over-fitting during the proxy/target calibration interval that ironically coincides with the industrial era (2).

**Data adaptation.** The three regional oak ring width subsets (NEF, NEG, SEG) and their CE mean compilation (ALL) were reduced in sample size, i.e., adapted, to test for possible effects of temporal replication changes on chronology behaviour (Fig. S1, S7; Table S1). The NEF data were reduced from originally 2880 to an adapted version of only 1025 series. The NEG data were reduced from 1785 to 833 series, the SEG data from 2619 to 779 series, and the CE mean dataset (ALL) from 7284 to 2637 series. This data adaptation procedure during which no new series were created, resulted in a more even distribution of individual series start and end dates throughout the past ~2500 years (Fig. S1), associated with a more homogeneous sample replication over time (Fig. S7C). Removal of samples from the original to the adapted datasets followed a random series selection process applied on highly-replicated pre-industrial periods from ~200 BC to AD 200, ~AD 400-900 and ~1000-1800, to roughly match sample size levels of the low-replication intervals from ~AD 200-400, as well as the 10th and 19th centuries (Fig. S1, Fig. S7C). The resulting adapted datasets and their subsequent oak chronologies (see below) allowed testing for biases introduced by temporal changes in sample size. The adaptation method further enabled us to consider possible uncertainties that might emerge from the integration of predominantly juvenile (fast growing) or mature/adult (slow growing) wood during specific periods of time.

**Growth-trend analysis.** Alignment of all individual raw ring width measurement series by their innermost ring, ideally representing the cambial age, facilitated the assessment of growth trends and levels (3). The resulting growth curves of the three regions, the so-called Regional
Curves (RCs) commonly describe trends of negative exponential shape (3), with very little differences between the original and adapted subsets (Fig. S2A). An assessment of the mean segment length (MSL) and average growth rate (AGR) of the raw measurement series indicated that shorter oak series, containing a greater fraction of juvenile wood, are characterized by overall higher growth rates, whereas longer series of more mature and adult wood contain generally lower growth rates (Fig. S2B). This commonly observed association between MSL and AGR in the raw ring width series emphasizes the need for tree-ring standardization (detrending) prior to any meaningful interpretation of externally forced variations (i.e., the putative climatic signal within the oak ring width data must be separated from the prevailing background noise). The observed similarities in AGR and consistency in the MSL/AGR association among the various regional subsets denoted their compatibility also with respect to growth rates and trends.

**High-frequency preservation.** To maintain inter-annual (high-frequency) variability from the three regional oak ring width subsets (NEF, NEG and SEG), the original and adapted subsets of raw measurement series were standardized, i.e., detrended using individual cubic smoothing splines with 50% frequency-response cut-off at 20 years (4). This detrending approach has been demonstrated to robustly preserve growth extremes while eliminating background noise on inter-decadal and longer time-scales (5). Annual ring width indices were herein calculated as ratios, but also as residuals after the application of a data adaptive power-transformation applied on the raw measurement series (6). A bi-weight robust mean was used to generate regional-scale subset chronologies of high-frequency variability. The various oak chronologies (ratios/original; residuals/original; ratios/adapted; residuals/adapted) per region were truncated at a minimum replication of five series and normalized over their individual length. To minimize biases due to replication and inter-series correlation changes, the normalized time-series were additionally corrected for artificial variance changes by calculating ratios from 31-year moving standard deviations (7), i.e., the normalized annual
chronology indices were divided by the corresponding values of their 31-year moving standard deviations (Fig. S3). The variance stabilized, regional-scale (high-frequency) chronologies are most suitable to detect inter-annual growth extremes likely caused by hydroclimatic anomalies. Nevertheless, they do not contain any information about longer-term changes in the prevailing climate system or their surrounding environment (5).

**Extreme-year validation.** We calculated moving 31-year correlation coefficients to assess the shared inter-annual (high-frequency) variability amongst the three regional subset chronologies (Fig. S4A). Correlation coefficients amongst the three time-series average at 0.44 over the common 210 BC to AD 1992 period. Higher agreement of $r = 0.52$ was found over the better-replicated last millennium. Reduced coherency amongst the three regional chronologies occurred from ~1800 onwards, the period during which the randomly updated recent data prevail. This recent coherency decline indicates that the proxy/target calibration results (see below) are likely conservative (2). The variance stabilized high-frequency chronologies (Fig. S4B) were further used to reconstruct the frequency and severity of annual growth extremes, which either occurred at the regional-scale, i.e., in two of the three regional chronologies, or at the sub-continental CE-scale, i.e., common to all three regional oak chronologies (Fig. S4C). Regional extremes in radial oak growth were defined as occurring when all four chronologies per region (ratios/original; residuals/original; ratios/adapted; residuals/adapted) exceeded their corresponding 1.5 standard deviation threshold (5). Sub-continental CE network extremes occurred when all three regional records shared an extreme year.

Scaled precipitation anomaly (1961-1990 mean subtracted) composites (8) for CE summers (June-August) were computed for the oak extremes back to AD 1500 (Fig. S5). In contrast to the classical compositing technique that uses the arithmetic mean and a t-test for the means, we applied a scaled mean and modified t-value (8). This method provides more robust results if the distribution is not ‘Gaussian’, in case of small samples or if there are
outliers. Scaled anomaly composites were calculated for the 12 most positive and the 16 most negative oak extremes back to AD 1500 using gridded precipitation reconstructions (9).

We extracted historical documentary evidence assembled and uploaded to the geo-database ‘Digital Atlas of Roman and Medieval Civilizations’ (DARMC; http://darmc.harvard.edu), to verify the climatic signal in the oak extremes further back in time, i.e., during the first half of the last millennium during which no gridded precipitation reconstructions exist (9). The early historical records in Latin, Middle Dutch, Middle French, and Middle German supplied a total of 87 different medieval records, usually eyewitneses from or near the regions represented by the oak network. The 87 different records offered 104 reports (since some records report on multiple years) on regional hydroclimatic conditions in 32 extreme years, resolved to the year or better, with 1-7 witnesses per year. A total of 88 out of 104 reports corroborate 30 out of 32 of the precipitation extremes preserved in our oak record between 1013 and 1504. Of the 16 contradictory reports, 13 occur for 9 years (AD 1221, 1270, 1309, 1350, 1353, 1368, 1434, 1464, 1487) whose precipitation extremes are in fact corroborated by 20 (of the 88) confirming historical witnesses. The contradictory reports for those 9 years may reflect local precipitation variation or simply ambiguity in those medieval written records. The only three reports that contradict the dendro-data in the absence of other, corroborating testimony, refer to two poorly documented years (AD 1044, 2 reports, and 1121, 1 report). Thus the medieval human witnesses independently confirm the skill of our method for reconstructing precipitation from the proxy data in a time and region where the anthropogenic impact on climate systems and mechanisms was minimal.

Three independently developed oak chronologies from Great Britain and Northern Germany (10), from Central Germany (11) and from Slovenia (12) were re-processed here, i.e., 20-year high-pass filtered, and then used for comparison with our new high-frequency oak data, both at the regional-scale and the CE network level (Fig. S6).
Tree-ring detrending. The three regional datasets including their original and adapted versions were horizontally split into recent and historical oak samples (2, 11), and processed using various detrending methods. This approach helped us not only to best maintain high-to-low-frequency hydroclimatic variability, but also to detect and account for biologically induced age trends (4, 6), population biases (3), and recent chronology characteristics (2, 11). Aligning the raw ring width measurement series by cambial age revealed common age trends in the NEF, NEG and SEG oak data. This alignment further demonstrated coherent relationships between MSL and AGR amongst the three regions (Fig. S2). Possible chronology biases related to changes in sample size and tree population have been discussed in the light of data adaptation (Fig. S1). Uncertainty associated with the most recent end of tree-ring chronologies that might emerge from exceptional changes in concentrations of atmospheric greenhouse-gases, levels of biospheric fertilization, the amount of forest management and degree of habitat opening (2), as well as end-effect problems in chronology behaviour (6), are most likely relevant for lower elevation temperate oak forest across CE (11), and thus justify horizontal data splitting into recent and historical subsets prior to their detrending (Table S1) (see also description below).

An array of eight different detrending methods using individual cubic smoothing splines with 50% frequency-response cut-off at 150 years (4) and alternatively using the Regional Curve Standardization (RCS) method (3), either based on ratios or residuals and utilizing the original or adapted datasets (ratios/original; residuals/original; ratios/adapted; residuals/adapted) was employed to test for possible effects on the resulting ring width chronologies. Note that correlation coefficients amongst the three regional (historical RCS/recent spline) chronologies either using the original or the adapted data (i.e., original versus adapted) average at 0.94 over the past two millennia. Correlation coefficients amongst the three regional (historical RCS/recent spline) chronologies either using ratios or residuals after power-transformation (i.e., ratios versus residuals) average at 0.93 over the past two
millennia. The different chronology versions (original versus adapted) after 150-year spline detrending show rather similar inter-annual to multi-decadal variations (Fig. S7). The corresponding Expressed Population Signal (EPS) values constantly range above the commonly applied threshold of 0.85 back to ~300 BC. No differences in mean EPS over the full period have been found between the original (0.97; 7284 series) and adapted (0.95; 2637 series) datasets (Fig. S7A). The EPS statistic, however, mainly verifies the high-frequency coherence of the oak measurements and potential changes in the wood source material that are common to the original and adapted datasets may still imply some non climatic noise. The 95% bootstrap confidence intervals of the original and adapted oak chronologies after 150-year spline detrending further reveal similar results (Fig. S7B), suggesting that the chronology behaviour is robust over time, and biases due to temporal changes in sample size are negligible. The chronologies (original/adapted) correlate at 0.93 with each other back to 400 BC, even though their replication substantially differs (Fig. S7C).

Some offset between the different tree-ring detrending and chronology development techniques was, however, found during the 20th century (Fig. S8). Those chronologies that were based on 150-year spline detrending contain less positive trends and best resembled the nearly trend-free gridded precipitation indices. Those chronologies that were developed with the RCS method and thus allow lower frequency variation to be preserved though indicate long-term growth increases over 1901-2006, possibly induced by effects of modern changes in forest management and cultivation, as well as further biases predominantly associated with the industrial period and recent end of tree-ring records including increased atmospheric greenhouse-gas, biospheric fertilization, sample replication, age-structure and chronology development. See (2) for a description of such effects and techniques to overcome them. We therefore considered the recent subsets of the three regions (NEG, NEF, SEG) after they were detrended with individual 150-year cubic smoothing splines combined with the RCS detrended historical series as the best method presently available to preserve hydroclimatic
information in an oak ring width chronologies. Similar techniques were applied in a comparable, but independent study from Central Germany (11).

Nevertheless, we are aware of the compromise this step implies: while minimizing biases that are most critical during the recent section of the millennial-long oak chronologies, the approach of horizontal splitting and using different detrending techniques at the same time limits the comparison of historical and recent information, and possible conclusions about the comparatively small hydroclimatic variations inferred for the recent period. Nevertheless, we believe this to be less critical, since our study mainly focuses on pre-industrial climate change. The historical subsets of the three regions were therefore detrended with the RCS method to maintain possible low-frequency information on time-scales exceeding those of the individual series lengths (13). This approach of horizontally different detrending, i.e., composite RCS for the historical samples and individual spline for the recent samples, was applied, as the recent chronology portion possibly contains biases (as discussed earlier), some of which are unique to the past ~150 years and the lower elevation temperate CE forests (2). Changes in forest management and habitat cultivation yielding increased forest productivity, via enhanced tree growth – which we believe to be most critical factors, however, possibly were less severe during earlier periods, but certainly were negligible for natural stands near the upper elevational treeline in the European Alps, for example, where human interventions played and play a less important role (see below). The recent spline and historical RCS chronologies were annually weighted by their sample replication and averaged. Note that even though the horizontal split approach may reduce some longer-term biases in the chronology (11, 14), it can possibly also diminish lower frequency information within the shorter recent subsets (13). This methodological constraint may complicate any straightforward comparison between modern industrial and earlier pre-industrial hydroclimatic variability.

**Precipitation reconstruction.** Due to non-significant differences between chronologies based on either the original or the adapted datasets, as well due to using ratios or residuals
after power-transformation for index calculation, it appeared unjustified to select a ‘single best solution’ and we therefore developed three regional-scale mean records considering information from the four minimally different chronologies. The three mean regional chronologies, including information from historical RCS chronologies and recent spline chronologies using ratios as well as residuals after power-transformation for indexing, were annually weighted by their individual sample replications and averaged to form a CE mean oak chronology. This procedure allowed us to firstly assess growth rates, trends and responses at the regional-scale, to secondly compare those parameters amongst the three regions as well against independent existing chronologies, and to finally compile regional evidence in a sub-continental network. It should further be noted that the region from where the NEF data are derived was under Roman occupation, that NEG region was never under Roman occupation and that SEG was at a border region.

We applied growth-climate response analyses of the new CE mean oak chronology using monthly resolved precipitation totals (mm/day) averaged over the 6-12° E and 48-52° N CE region (16) (Fig. S9). Previous year precipitation showed no impact on oak growth, whereas monthly totals of April, May and June revealed positive correlations (Fig. S9C). April-June (AMJ) precipitation totals revealed the highest positive correlation, significant at the 99.9% confidence limit (1901-1980), and were subsequently used for reconstruction purposes. To avoid regression-based variance reduction in the proxy reconstruction model (15), the CE mean oak chronology was scaled (1901-1980) against April-June (AMJ) precipitation totals (mm) averaged over the 6-12° E and 48-52° N CE region (16). This ‘composite plus scaling’ procedure, i.e., the adjustment of proxy mean and variance, is the simplest amongst various calibration techniques but is perhaps also least prone to variance underestimation (for a more detailed discussion see 15, 17). The relationship between radial oak growth and AMJ precipitation was tested for temporal and spatial stability (Fig. S9). Field correlation analyses between the oak chronology and gridded precipitation indices (16) were performed for the
European sector (Fig. S9D). The spatial signature of the oak proxy was compared with the idealized spatial correlation field obtained from the mean of three instrumental precipitation target records that best represent the location of the tree oak sample regions (i.e., stations from Nancy, Regensburg and Potsdam). Uncertainty bars of the final reconstruction reflect the +/- 1 root mean square error (RMSE) computed from the calibration period.

**Temperature reconstruction.** A total of 1546 ring width series from high-elevation conifers sampled in the Austrian Alps, mainly Tyrol and adjacent areas was aggregated. This compilation describes an updated subset of a near Holocene-long tree-ring network of subfossil wood, historical timber and recent trees (18). The dataset contains the two dominant Alpine treeline species Stone pine (*Pinus cembra*) and European larch (*Larix decidua*), with their radial growth being dominated by summer temperature variability (14). A total of 1089 series represents pine trees and 457 series represent larch trees, either containing the pith or allowing pith-offsets to be estimated. The combined dataset is characterized by sufficient sample size, i.e., a mean of 109 series per year over the past 2500 years (Fig. S10). Constant replication of >100 series is given from AD 1249 onwards, whereas a sample size of >30 series reaches back to AD 134.

Data were horizontally separated, RCS detrended, weighted averaged and scaled against homogenized JJA temperature means of the great Alpine region (19,20) and the 1864-2006 period (Fig. S11). A similar approach and interval has been previously demonstrated to allow variations in Alpine summer temperature to be robustly reconstructed for the past 1250 years (21). Field correlation analysis (22) between the new Alpine conifer chronology and gridded JJA temperature indices (16) was performed over the European sector (Fig. S12). Uncertainty bars refer to the +/- 1 RMSE derived from the calibration period.
Figures and tables

**Fig. S1.** Series replication of the three historical oak datasets (NEF, SEG, NEG), using either their original or adaptation subsets (see table S1 and figure S7 for details).
Fig. S2. (A) Regional Curves of the age-aligned TRW measurement series and their replication using the original and adapted regional sub-sets (see table S1 for details). (B) Relationship between average growth rate (AGR) and mean segment length (MSL) of the 7284 oak samples geographically divided into the three regional sub-sets (Northeast France = NEF, Southeast Germany = SEG, Northeast Germany = NEG).

Fig. S3. Moving 31-year standard deviations of the three regional high-frequency chronologies (after 20-year spline detrending with and without power-transformation and using the original and adapted subsets) before (black) and after (color) variance stabilization.
Fig. S4. (A) Moving 31-year correlations and Rbar (black) between (B) the regional sub-set chronologies (after 20-year spline detrending and variance stabilization), and (C) positive and negative extremes amongst the entire network (black) and the regional sub-sets.

Fig. S5. Scaled anomaly composites (8) of summer precipitation (9) computed for (A) the 16 negative and (B) the 12 positive oak extremes between AD 1500 and 2000.
Blue and red colours refer to wet and dry conditions. The green shadings (Right) indicate the 95% confidence levels.

**Fig. S6.** Positive and negative oak extremes that either occurred over the entire network (black), or were restricted to the three regional subsets NEF/SEG (red), NEG/NEF (green) and SEG/NEG (blue) back to 1000. Corresponding colour-arrows refer to independent oak data from Central Germany (11; 1063-1990; B), Slovenia (12; 1497-2003; C) and Great Britain (10; 0-1996; K) that show similar growth extremes.

**Fig. S7.** (A) The Expressed Population Signal (EPS) of the original (blue; 7284 series) and adapted (red; 2637 series) oak chronologies, computed over 50-year periods and shifted by 25 years back in time. (B) The 95% bootstrap confidence intervals of the original (blue) and adapted (red) oak chronology after 150-year spline detrending. The time-series are 20-year low-pass filtered. (C) Series replication of the original (blue) and adapted (red) chronologies.
Fig. S8. (A) Precipitation anomalies (mm/day with respect to 1961-1990; CRUTS3) averaged over the 6-12° E and 48-52° N CE region, compared with different chronology versions of the three regional subsets: (B) NEF, (C) SEG and (D) NEG. The light colors refer to four RCS chronologies (ratios/original; residuals/original; ratios/adapted; residuals/adapted) that allow trend biases to occur, whereas the bright colors refer to the corresponding four chronology versions after 150-year spline chronologies that prevent possible biases during the recent time-series ends.
Fig. S9. (A) Moving 31-year correlations between AMJ precipitation (averaged over 6-12° E and 48-52° N) and the mean CE oak chronology. (B) Measured (blue) and reconstructed (green) AMJ precipitation anomalies (mm/day with respect to 1961-1990; CRUTS3) after scaling over the 1901-1980 period. (C) Correlations between monthly precipitation totals and the oak chronology (1901-1980), and (D) spatial spearman correlations of the oak chronology against gridded (0.5°x0.5°) AMJ precipitation anomalies. The right map shows spatial correlations of the mean of three instrumental stations (Nancy, Regensburg, Potsdam) best representing the oak sub-regions against gridded (0.5°x0.5°) AMJ precipitation anomalies.

Fig. S10. Temporal distribution of the 1546 Alpine conifer ring width series from the Austrian Alps, classified into recent, historical and sub-fossil material, which have been used to reconstruct JJA temperature variability.
Fig. S11. (A) Moving 31-year correlations between the measured JJA temperatures of the Greater Alpine Region (20,21) and the Alpine conifer chronology. (B) Measured (red) and reconstructed (green) JJA temperature (°C) anomalies (wrt. 1961-1990) after scaling over the reliable 1864-2003 period of proxy/target overlap. The smoothed lines are 20-year low-pass filtered, corresponding to correlations in brackets.
Fig. S12. Spatial signature of the new CE JJA temperature reconstruction (black star) expressed as field correlations against gridded 0.5°x0.5° JJA temperatures (16) computed over the 1901-2003 period.

Table S1. Summary information of the oak ring width data used in this study. The inventory of the entire dataset (7284 series; ALL) is further divided into regional (NEF, NEG, SEG) and horizontal (recent and historic) subsets, considering the original (black) and adapted (grey) data versions. MSL = Mean segment length; AGR = Average growth rate; AC1 = First order autocorrelation; Start/End >5 = Chronology start and end dates with a sample replication >5 series; NEF = Northeast France; NEG = Northeast Germany; SEG = Southeast Germany.

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Table S2. Inventory of the annual oak growth extremes detected at the overall CE and the regional-scales over the entire time-series spans. B covers the AD 1063-1990 (11); C covers the AD 1497-2003 (12); K covers the AD 0-1996 (10); * have been used for the anomaly composites.
References


