



Supporting Online Material for
Propagation of Slow Slip Leading Up to the 2011 M_w 9.0 Tohoku-Oki Earthquake

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This PDF file includes:

Materials and Methods
Fig. S1
Table S1
Full Reference List

Materials and Methods

Methods of a matched filter technique

The analysis procedure of a matched filter technique follows that of *Shelly et al. (15)* and *Peng and Zhao (16)*. We here describe the procedure briefly. To detect missing foreshock events between 13 February and 11 March in 2011, we analyzed the continuous three-component velocity seismograms recorded at 14 stations along the Pacific coast for the same period (Fig. 1), operated by the National Research Institute for Earth Science and Disaster Prevention (NIED) and Tohoku University. As template events, we selected all 333 earthquakes listed in the JMA catalog between 13 February and 11 March (the M_w 9.0 mainshock origin).

For each template event at each station, we used 6 seconds of the waveform, starting 3 second before the S-wave arrival time computed using a one-dimensional velocity model proposed by the JMA. After down-sampling of the data from 100 Hz sampling frequency to 20 Hz, we applied a two-way 1–6 Hz Butterworth filter to the data. We calculated the correlation coefficient as a function of time, shifting the window in increments of 0.05 seconds through the continuous waveforms (see Fig. S1). At each point, the correlation coefficient between a template event waveform and a target waveform was calculated on every component at every station, and was averaged throughout. The same procedure was repeated for each template event.

We used the mean correlation coefficient as our detection statistic. To define a detection threshold we used the standard deviation (sigma) of the mean correlation coefficients for each event. We set a conservative threshold for a positive detection at 8 times sigma. This level was chosen to suppress spurious detections while retaining as much legitimate detection as possible. When the mean correlation peak exceeded this threshold level, it was labeled as a positive detection in a statistically significant way. As the detected events have similar waveforms at multiple stations as compared to the template event, their hypocenter locations must be close or identical (*15, 16*).

We assigned the location of the detected event to that of the template event. For multiple detections in each 3 s window, we used the location of the template event with the highest mean correlation coefficient value. We assigned the time associated with the highest mean correlation coefficient value to its origin time of the detected event by subtracting the computed S-wave arrival time. Using relative distances between the multiple detected events, we estimated the average relative uncertainty in the detected event locations to be approximately 4 km. Finally, we computed the magnitude of the detected event based on the median value of the maximum amplitude ratios for all channels between the template and detected events, assuming that a tenfold increase in amplitude corresponds to one unit increase in magnitude (*16*).

Detection of repeating earthquakes

To identify repeating earthquakes in the JMA catalog, we used waveform data of earthquakes of M2.0 or larger listed in the catalog for the period from January 2002 to 11 March in 2011, in the studied area. We picked up earthquake pairs whose epicenter separations are up to 20 km, and calculated cross-correlation coefficients using 1–4, 2–8, and 4–16 Hz passband filtered

waveforms. The time window was from P-wave onset to three seconds after direct S-wave arrival. We selected as a repeating earthquake pair, those whose cross-correlation coefficients, for the frequency range roughly corresponding to the corner-frequency of each event, were greater than 0.95 at two or more stations (23).

The detection criteria for the newly detected events (green stars in Figs. 2 and 3), which were found to resemble the repeating earthquakes in the JMA catalog, are the same as for the other events.

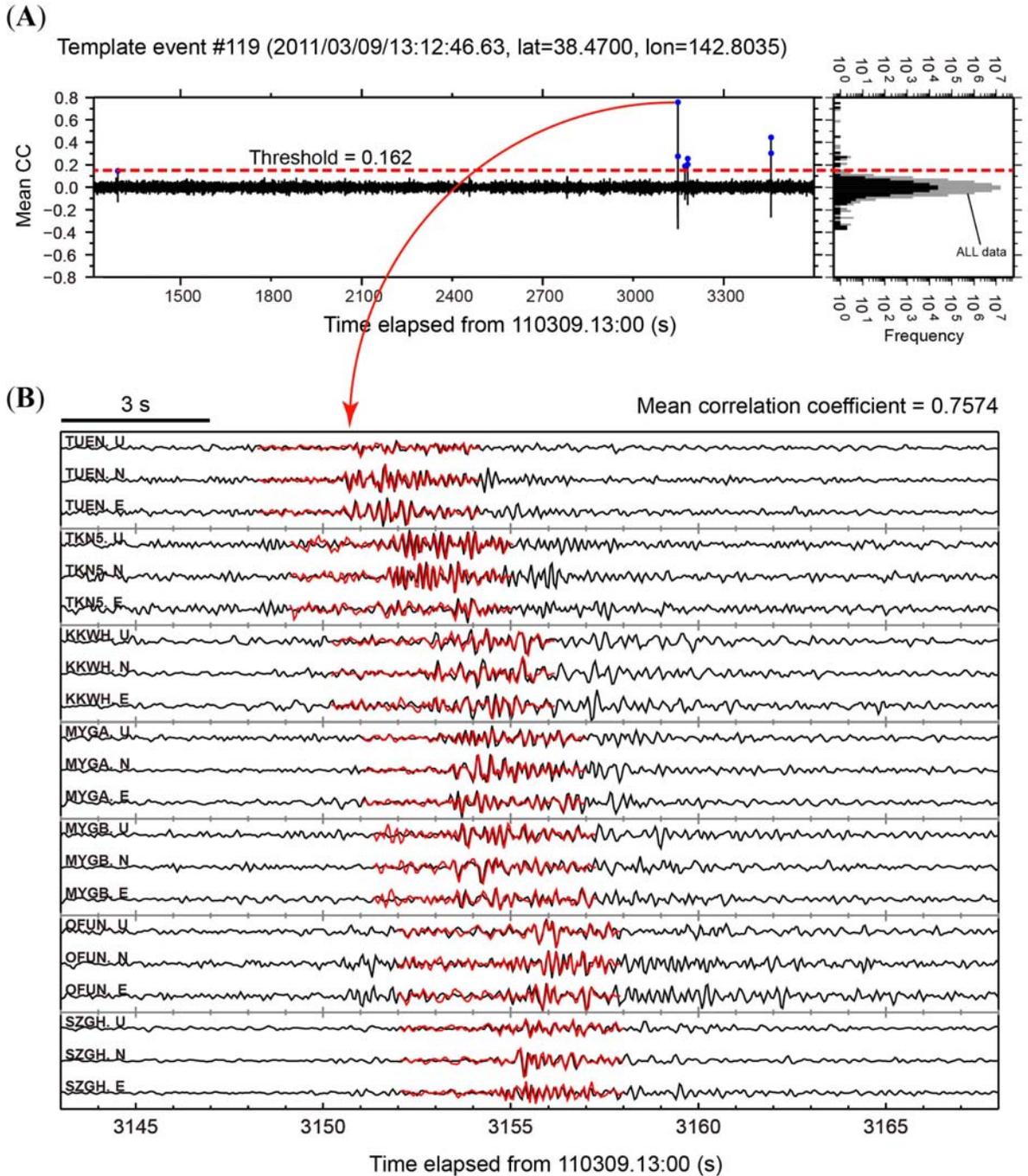


Figure S1 Mean correlation function and detection example. (A) Mean correlation function for one template event #119 against elapsed time from March 09, 2011, 13:00. Histogram of mean correlation values within this time window and entire period are shown to the right as black and gray bars. (B) Waveforms at the time detection for template event #119. Continuous earthquake waveforms are shown in black and template event waveforms in red for each component of 6 seismic stations arranged according to the epicentral distance. The station names and components are given to the left of each trace. Waveforms are bandpass filtered between 1 and 6 Hz, and template event amplitudes are scaled to match the continuous data.

Table S1. List of the time (T_h) required for afterslip to release half of the mainshock moment following major M7–8 plate-interface earthquakes that took place along the Japan and Kuril trenches. We used analysis results based on GEONET, a dense GPS geodetic network.

Event date	Longitude	Latitude	Depth(km)	M_w	T_h (months)	References
1994/12/28	143.7450	40.4300	0.00	7.6	2.23	(26)
2003/09/26	144.0785	41.7785	45.07	8.0	10.24	(35)
2004/11/29	145.2755	42.9460	48.17	7.0	0.80	(36)
2005/08/16	142.2778	38.1495	42.04	7.1	2.04–3.53	(32)*, (37)
2008/07/19	142.2645	37.5208	31.55	7.0	1.64	(32)*
2010/03/14	141.8180	37.7242	39.75	6.7	1.27	(32)*
2011/03/09	143.2798	38.3285	8.28	7.3	0.07	(29) This study

*We estimated the T_h values using average decay time determined for each station by *Suito et al.* (2011) (32), assuming an exponential decay of the afterslip moment.

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